

Formulas/Constants: PHYS/OCEA 4412/5412 A Atmospheric Dynamics II

$$\Gamma_d = 9.8 \text{ }^\circ\text{C/km (dry adiabatic lapse rate)}$$

$$\gamma = c_p/c_v = 1.4$$

$$R = 287.05 \text{ J/(kg K)}$$

$$c_v = 717.5 \text{ J/(kg K)}$$

$$c_p = 1004.5 \text{ J/(kg K)}$$

$$g_0 = 9.81 \text{ m/s}^2$$

$$f = 2\Omega \sin \phi_0$$

$$\Omega = 7.292 \times 10^{-5} \text{ rad/sec}$$

$$p_0 = 1013.25 \text{ hPa}$$

$$M_d = 28.94 \text{ g/mole (Mean molecular weight of dry air)}$$

$$\nu = 1.46 \times 10^{-5} \text{ m}^2/\text{sec (sea level)}$$

$$\Gamma_d = 9.8 \text{ }^\circ\text{C/km}$$

$$a = 6370 \text{ km (radius of the earth)}$$

$$\theta = T(p_s/p)^{\frac{(\gamma-1)}{\gamma}} = T(p_s/p)^{0.286} = T(p_s/p)^{R/c_p}, \text{ where } p_s = 10^5 \text{ Pa}$$

$$\text{Ideal Gas Law : } p = \rho RT$$

$$\text{specific volume: } \alpha = 1/\rho$$

$$\text{Pressure Gradient Acceleration : } \mathbf{F}/m = -\frac{1}{\rho} \nabla p = -\nabla_p \Phi$$

$$\text{Definition of Geopotential : } d\phi = gdz$$

$$\text{Hydrostatic Relationship : } dp/dz = -\rho g$$

$$\text{Thickness Relationship : } dz = -\frac{R\langle T \rangle}{g} d \ln p$$

$$\text{Total vs Local Derivative : } d/dt = \partial/\partial t + \mathbf{U} \cdot \nabla$$

$$\text{Momentum Equation : } d\mathbf{U}/dt = -2\boldsymbol{\Omega} \times \mathbf{U} - (1/\rho)\nabla p + \mathbf{g} + \mathbf{F}_r$$

$$\text{Geostrophic wind : } \mathbf{V}_g = \mathbf{k} \times (1/(\rho f))\nabla p$$

$$u_g = -\frac{1}{\rho f} \frac{\partial p}{\partial y}$$

$$v_g = \frac{1}{\rho f} \frac{\partial p}{\partial x}$$

$$\text{Momentum Equation p surface : } d\mathbf{V}/dt + f\mathbf{k} \times \mathbf{V} = -\nabla_p \phi$$

$$\text{pressure velocity : } \omega = dp/dt$$

$$\text{Geostrophic wind on pressure surface : } \mathbf{V}_g = \mathbf{k} \times (1/f)\nabla_p \phi$$

In component form:

$$u_g = -\frac{1}{f} \frac{\partial \phi}{\partial y}$$

$$v_g = \frac{1}{f} \frac{\partial \phi}{\partial x}$$

$$\text{Continuity equation : } \frac{\partial \rho}{\partial t} + \nabla \cdot (\rho U) = 0$$

$$\text{Continuity equation : } \frac{1}{\rho} \frac{d\rho}{dt} + \nabla \cdot U = 0$$

$$\text{Thermodynamic equation : } c_v \frac{dT}{dt} + p \frac{d\alpha}{dt} = J$$

$$\text{Thermodynamic equation : } c_p \frac{dT}{dt} - \alpha \frac{dp}{dt} = \dot{Q}$$

$$\frac{T}{\theta} \frac{\partial \theta}{\partial z} = \frac{\partial T}{\partial z} + \frac{g}{c_p} = -\Gamma + \Gamma_d$$

$$v_g = \frac{1}{f} \frac{\partial \phi}{\partial x}$$

$$u_g = -\frac{1}{f} \frac{\partial \phi}{\partial y}$$

$$\text{Continuity equation in pressure coordinates : } \left(\frac{\partial u}{\partial x} + \frac{\partial v}{\partial y} \right)_p + \frac{\partial \omega}{\partial p} = 0$$

$$\text{Thermodynamic equation in pressure coordinates : } \left(\frac{\partial T}{\partial t} + u \frac{\partial T}{\partial x} + v \frac{\partial T}{\partial y} \right) - \omega S_p = \frac{J}{c_p}$$

$$\text{Static stability : } S_p = -\frac{T}{\theta} \frac{\partial \theta}{\partial p} = -\frac{\partial T}{\partial p} + \frac{RT}{c_p p} = \frac{\Gamma_d - \Gamma}{\rho g}$$

$$\text{Thermal Wind : } p \frac{\partial v_g}{\partial p} = -\frac{R}{f} \left(\frac{\partial T}{\partial x} \right)_p$$

$$p \frac{\partial u_g}{\partial p} = \frac{R}{f} \left(\frac{\partial T}{\partial y} \right)_p$$

Other forms involving Φ

$$f_0 \frac{\partial v_g}{\partial p} = \frac{\partial^2 \phi}{\partial x \partial p}$$

$$f_0 \frac{\partial u_g}{\partial p} = -\frac{\partial^2 \phi}{\partial y \partial p}$$

Absolute vector vorticity : $\omega_{\mathbf{a}} = \nabla \times \mathbf{U}_{\mathbf{a}}$

Relative vector vorticity : $\omega = \nabla \times \mathbf{U}$

Relative scalar vorticity : $\xi = \frac{\partial v}{\partial x} - \frac{\partial u}{\partial y}$

Total (Absolute) vorticity : $\eta = \xi + f$

Thermal Wind in Vector Form : $\mathbf{V}_T = \frac{R}{f} \mathbf{k} \times \nabla \langle T \rangle \ln(p_0/p_1)$

$$u_T = -\frac{R}{f} \left(\frac{\partial \langle T \rangle}{\partial y} \right)_p \ln\left(\frac{p_0}{p_1}\right)$$

$$v_T = \frac{R}{f} \left(\frac{\partial \langle T \rangle}{\partial x} \right)_p \ln\left(\frac{p_0}{p_1}\right)$$

$$u_T = -\frac{1}{f} \frac{\partial(\phi_1 - \phi_0)}{\partial y}$$

$$v_T = \frac{1}{f} \frac{\partial(\phi_1 - \phi_0)}{\partial x}$$

$$\omega \approx -\rho g w$$

Kinematic Estimate of ω : $\omega(p) = \omega(p_s) - \int_{p_s}^p \left(\frac{\partial u}{\partial x} + \frac{\partial v}{\partial y} \right) dp$

Surface Pressure Tendency : $\frac{\partial p_s}{\partial t} \approx - \int_0^{p_s} (\nabla \cdot \mathbf{V}) dp$

Relative scalar vorticity : $\xi = \frac{\partial v}{\partial x} - \frac{\partial u}{\partial y}$

Total (Absolute) vorticity : $\eta = \xi + f$

$$\text{Ertel Potential Vorticity : } P = (\xi_\theta + f) \left(-g \frac{\partial \theta}{\partial p} \right)$$

$$\xi_g = \frac{1}{f_0} \nabla^2 \Phi$$

$$\text{QG Vorticity Equation(5.45) : } \frac{\partial \xi_g}{\partial t} = -\mathbf{V}_g \cdot \nabla (\xi_g + f) + f_0 \frac{\partial \omega}{\partial p}$$

Adiabatic QG Energy Equation (5.48) :

$$\frac{\partial}{\partial t} \left(-\frac{\partial \Phi}{\partial p} \right) = -\mathbf{V}_g \cdot \nabla \left(\frac{\partial \Phi}{\partial p} \right) + \sigma \omega$$

$$\sigma = \frac{RS_p}{p} = -\alpha \frac{\partial \ln \theta}{\partial p}$$

Vorticity Equation Pressure Surfaces :

$$\frac{\partial \xi}{\partial t} = -\mathbf{V} \cdot \nabla (\xi + f) - \omega \frac{\partial \xi}{\partial p} - (\xi + f) \nabla \cdot \mathbf{V} + \mathbf{k} \cdot \left(\frac{\partial \mathbf{v}}{\partial p} \times \nabla \omega \right)$$

$$\mathbf{V}_{ag} = \frac{\mathbf{k}}{f} \times \frac{d\mathbf{V}}{dt}$$

$$\frac{d\mathbf{V}}{dt} - \left(\frac{d\mathbf{V}}{dt} \right)_0 = (\mathbf{V}_s \cdot \nabla) \mathbf{V}_0 + \frac{d\mathbf{V}_s}{dt}$$

$$\mathbf{V}_{ag,U} - \mathbf{V}_{ag,L} = \frac{\mathbf{k}}{f} \times \left[(\mathbf{V}_s \cdot \nabla) \mathbf{V}_0 + \frac{d\mathbf{V}_s}{dt} \right]$$

$$\mathbf{V}_{ag,I} = -\frac{1}{f^2} \nabla \frac{\partial \Phi}{\partial t}$$

$$\nabla \cdot \mathbf{V}_{isal} = -\frac{1}{f^2} \nabla^2 \frac{\partial \Phi}{\partial t}$$

$$\mathbf{V}_{IA} = \frac{\mathbf{k}}{f} \times \mathbf{V}_g \cdot \nabla \mathbf{V}_g$$

$$f_0 (\nabla \cdot \mathbf{V} - \nabla \cdot \mathbf{V}_0) = -\mathbf{V}' \cdot \nabla (\xi_{g0} + \xi_g + f)$$

$$\text{QG } \omega \text{ Equation : } \sigma \left(\nabla^2 + \frac{f_0^2}{\sigma} \frac{\partial^2}{\partial p^2} \right) \omega = f_0 \frac{\partial}{\partial p} [\mathbf{V}_g \cdot \nabla (\xi_g + f)] + \nabla^2 \left[\mathbf{V}_g \cdot \nabla \left(-\frac{\partial \Phi}{\partial p} \right) \right]$$

$$\mathbf{Q} = (Q_1, Q_2) = \left(-\frac{R}{p} \frac{\partial \mathbf{V}_g}{\partial x} \cdot \nabla T, -\frac{R}{p} \frac{\partial \mathbf{V}_g}{\partial y} \cdot \nabla T \right)$$

$$\sigma \left(\nabla^2 + \frac{f_0^2}{\sigma} \frac{\partial^2}{\partial p^2} \right) \omega = -2 \nabla \cdot \mathbf{Q}$$

In a natural coordinate system where the isotherms are parallel to the s direction with the warm air on the right:

$$\mathbf{Q} = -\frac{R}{p} \left| \frac{\partial T}{\partial n} \right| \left(\mathbf{k} \times \frac{\partial \mathbf{V}_g}{\partial s} \right)$$

$$\mathbf{Q} = f \gamma \frac{d}{dt_g} \nabla \theta$$

$$\mathfrak{S} = \frac{d|\nabla_p \theta|}{dt}$$

$$\mathfrak{S}_x = \frac{d}{dt} \left(\frac{\partial \theta}{\partial x} \right)$$

$$\mathfrak{S}_x = \frac{d}{dx} \left(\frac{d\theta}{dt} \right) - \frac{\partial u}{\partial x} \frac{\partial \theta}{\partial x} - \frac{\partial v}{\partial x} \frac{\partial \theta}{\partial y} - \frac{\partial \omega}{\partial x} \frac{\partial \theta}{\partial p}$$

$$\mathfrak{S}_{2D} = \frac{-1}{2|\nabla \theta|} \left[D \left(\left(\frac{\partial \theta}{\partial x} \right)^2 + \left(\frac{\partial \theta}{\partial y} \right)^2 \right) + F_1 \left(\left(\frac{\partial \theta}{\partial x} \right)^2 - \left(\frac{\partial \theta}{\partial y} \right)^2 \right) + 2F_2 \left(\frac{\partial \theta}{\partial x} \frac{\partial \theta}{\partial y} \right) \right]$$

$$\mathfrak{S}_{2D} = \frac{|\nabla \theta|}{2} (F \cos 2\beta - D)$$

$$D = \frac{\partial u}{\partial x} + \frac{\partial v}{\partial y}$$

$$F_1 = \frac{\partial u}{\partial x} - \frac{\partial v}{\partial y}$$

$$F_2 = \frac{\partial v}{\partial x} + \frac{\partial u}{\partial y}$$

$$F = (F_1^2 + F_2^2)^{\frac{1}{2}}$$

β : angle between isentropes and local dilation axis

$0^\circ < \beta < 45^\circ$: F promotes frontogenesis

$45^\circ < \beta < 90^\circ$: F promotes frontolysis

$$\mathfrak{S}_{2D} = \left(\frac{1}{f\gamma} \right) \frac{\mathbf{Q} \cdot \nabla \theta}{|\nabla \theta|}$$

$$M_g = u_g - fy$$

Inertial Instability (with f constant): $\partial M_g / \partial y > 0$

QG Geopotential Tendency Equation $\left(\chi = \frac{\partial \Phi}{\partial t} \right)$:

$$\left(\nabla^2 + \frac{f_0^2}{\sigma} \frac{\partial^2}{\partial p^2} \right) \chi = -f_0 \mathbf{V}_g \cdot \nabla \left(\frac{1}{f_0} \nabla^2 \Phi + f \right) - \frac{f_0^2}{\sigma} \frac{\partial}{\partial p} \left[\mathbf{V}_g \cdot \nabla \left(\frac{\partial \Phi}{\partial p} \right) \right]$$

$$T = -\frac{p}{R} \frac{\partial \Phi}{\partial p}$$

Penetration depth of PV anomaly: $H = fL/N$

$$\frac{dPV}{dt} \approx -g(\xi_\theta + f) \frac{\partial \dot{\theta}}{\partial p}$$

$$c = \bar{u} - \beta/k^2$$

$$\beta = \frac{\partial f}{\partial y}$$